

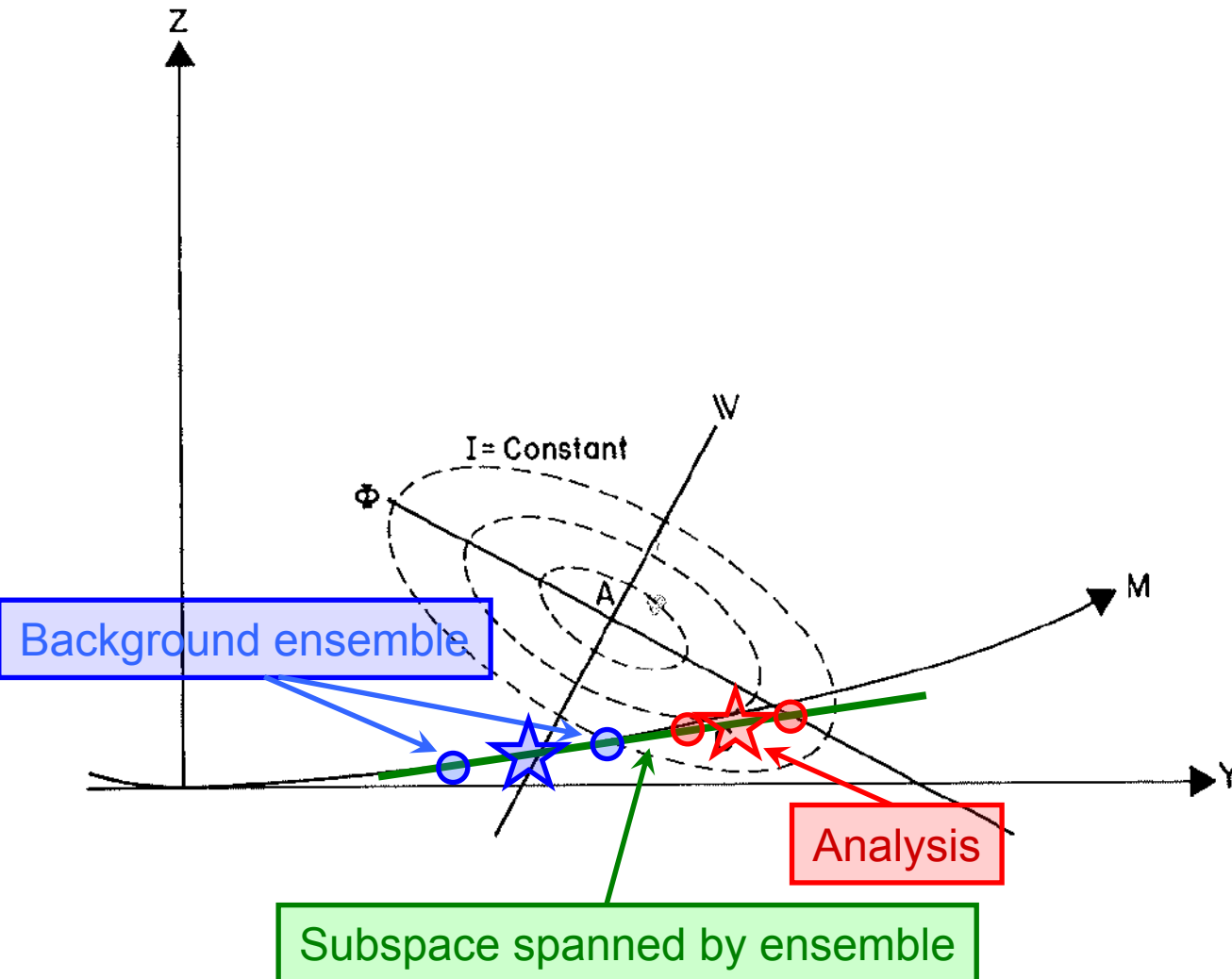
Balance in the EnKF

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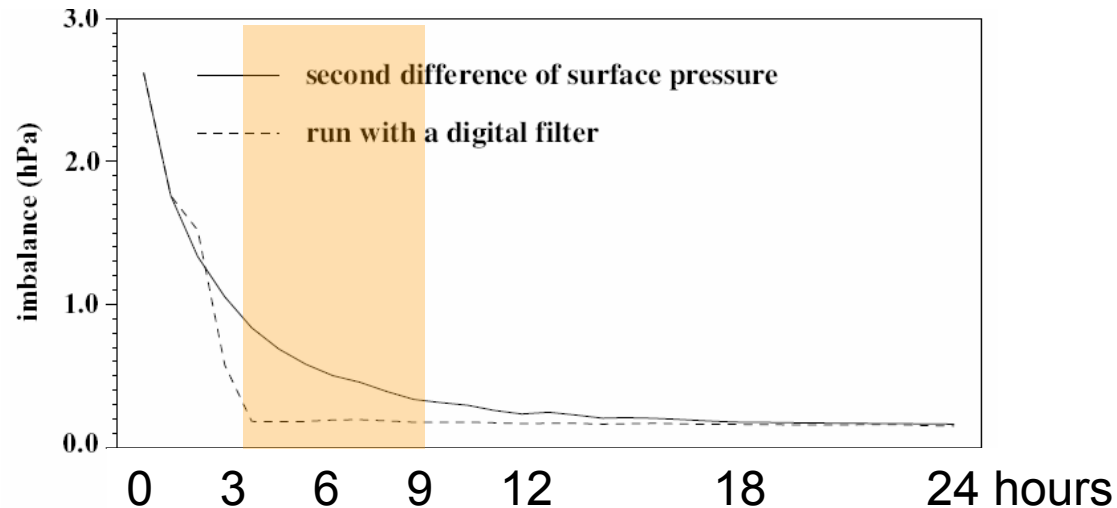
A partnership between the Australian Bureau of Meteorology and CSIRO
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EnKF and Slow Manifold



- Analysis is from the subspace spanned by the background.
- If the background is on the manifold, and the manifold curvature is small (relative to the analysis increment), then the analysis will be close to balanced.

Canadian experience

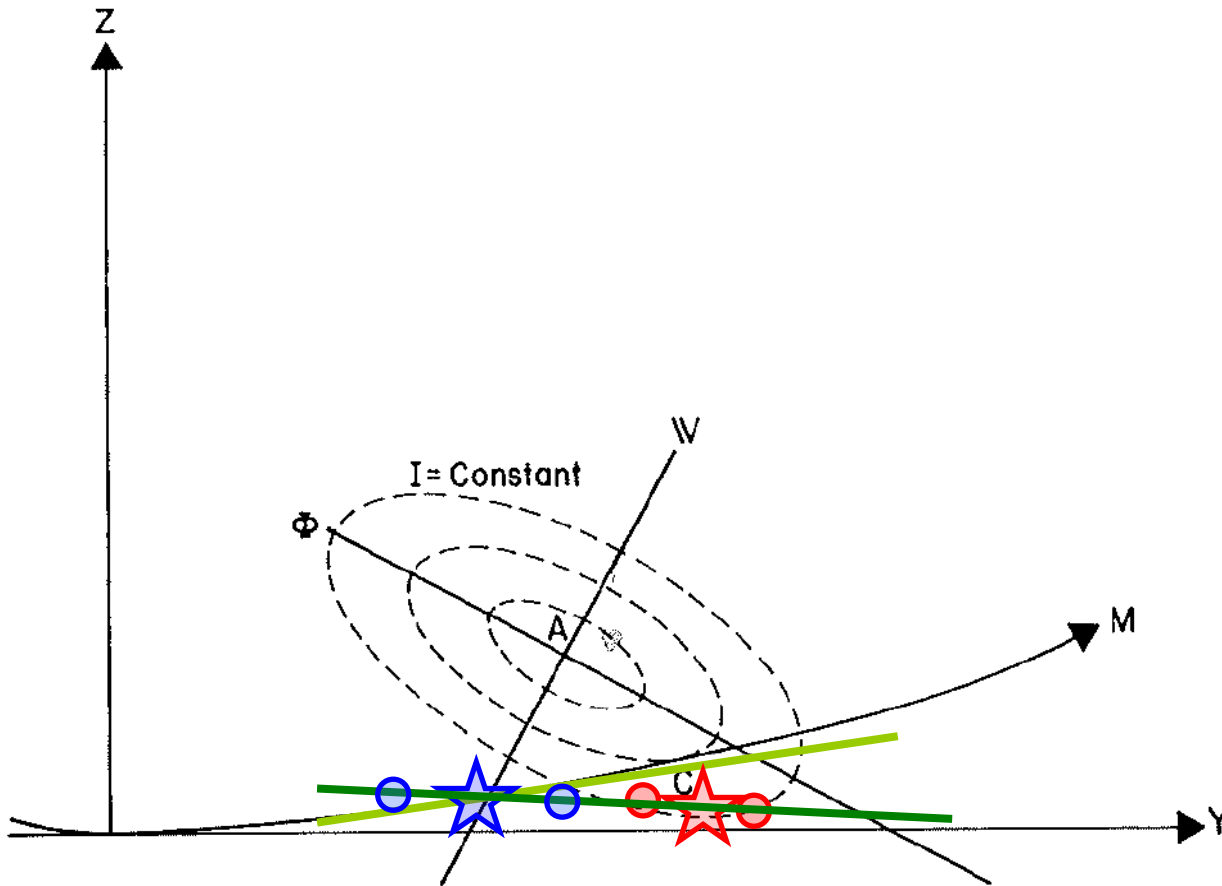


- An initialisation step is necessary as gravity waves in 3 to 9-hr forecast affect the ensemble covariances.
- Even though their model strongly damps gravity waves. (Houtekamer and Mitchell 2005)

Causes of imbalance in an EnKF

- Linear operations when balances are nonlinear
 - Analysis is in subspace spanned by ensemble
 - Covariance inflation (small)
- Background ensemble
- Approximations
 - Sampling error
 - Covariance localisation
- Mixed-mode (i.e. real) observations
- Introduced imbalance
 - Model error term
 - Blend with 3d-var covariances (“hybrid schemes”)
 - Initialisation balance \neq model balance (hopefully small)
- Can be hard to diagnose cause in a full system (Houtekamer and Mitchell 2005)

EnKF and Slow Manifold



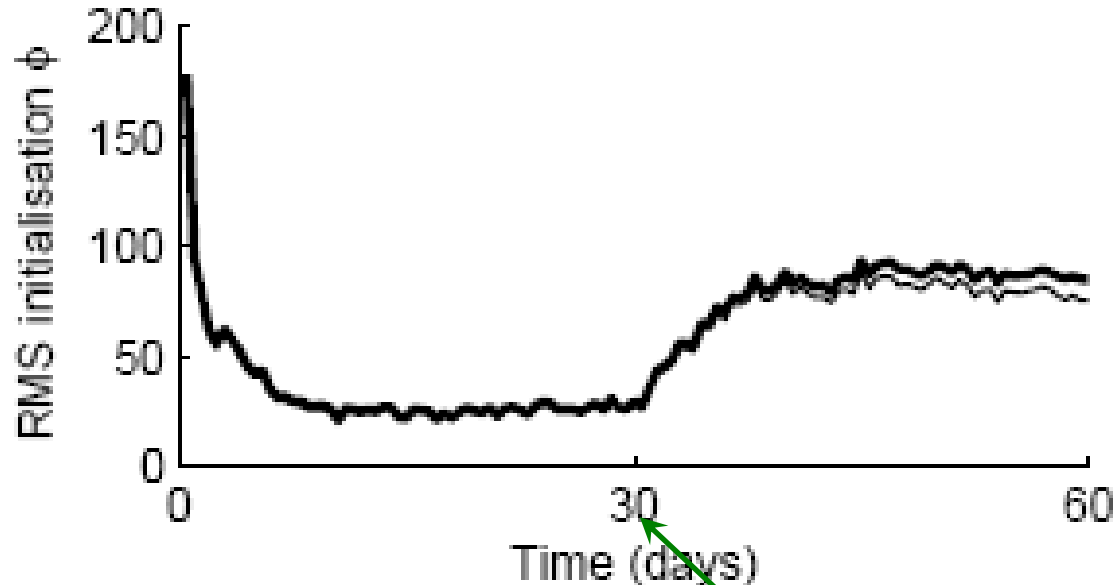
- Analysis is from the subspace spanned by the background.
- If the background is unbalanced, then the analysis probably will be also.

Positive feedback on imbalance

Less Balanced

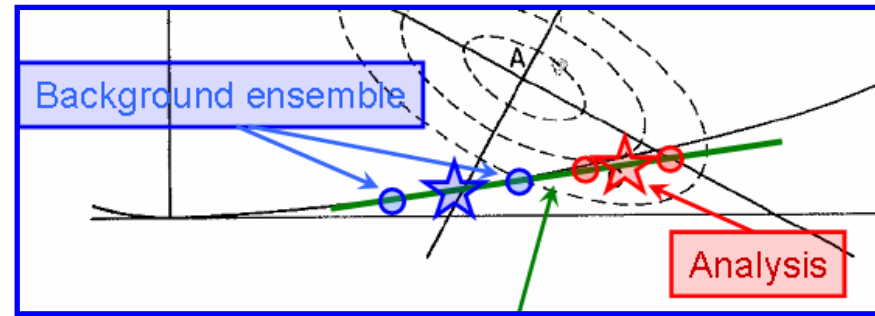


More Balanced



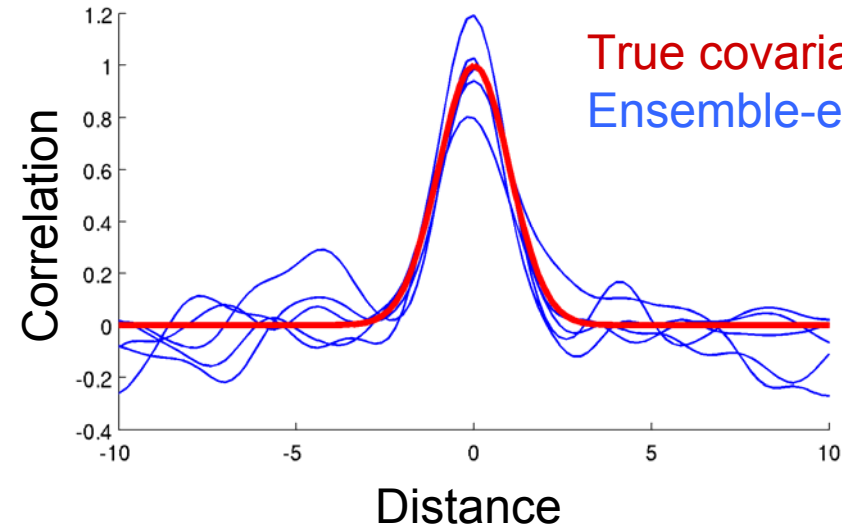
- Run EnKF with initialisation, then **turn NMI off.**
- Imbalance grows.
- Positive feedback: gravity waves describe increasingly more of the background variance (Lorenz 2003).

The problem with:

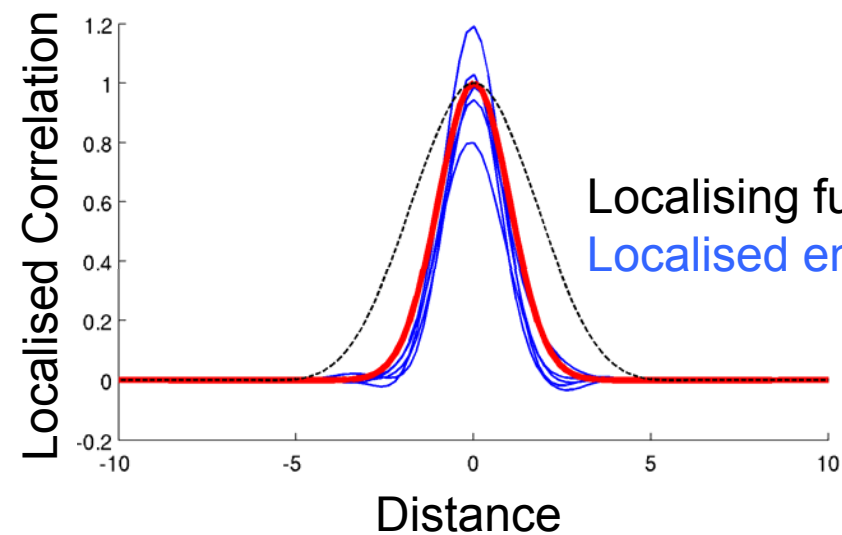


- The model has dimension $\sim 10^6$.
- The manifold has dimension $\sim 10^4$.
- The ensemble has dimension $\sim 10^2$.
- Hence the closest point to the observations on the ensemble-subspace will almost certainly be a long way from the closest point on the manifold.
 - i.e. The analysis will fit the observations poorly (e.g. Lorenc 2003).
- Thus localisation is necessary.

Covariance Localisation



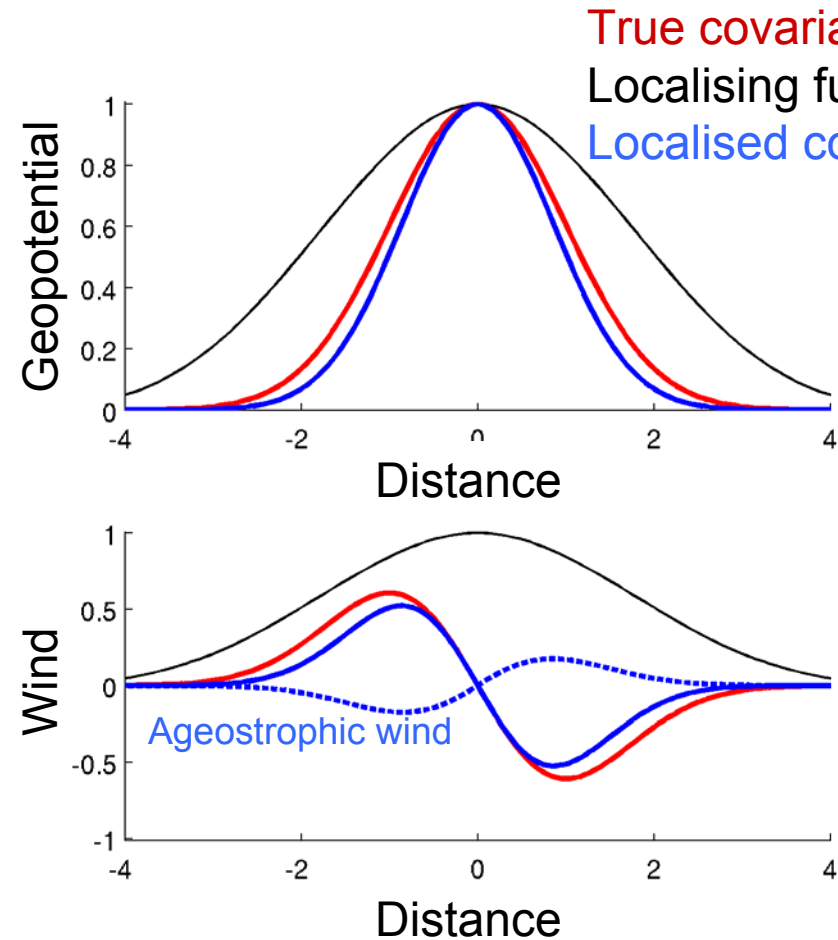
- Localisation eliminates the effect of sampling error at large distance (but not small).
- See Tom Hamill's talk.



Localisation and balance

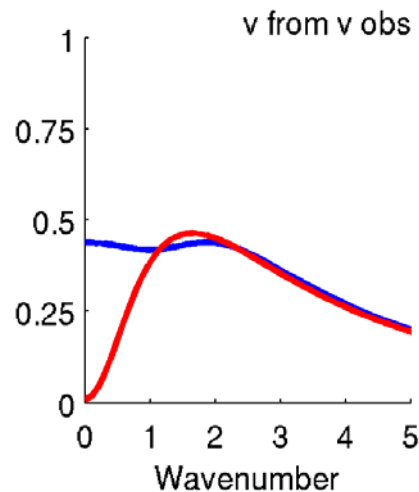
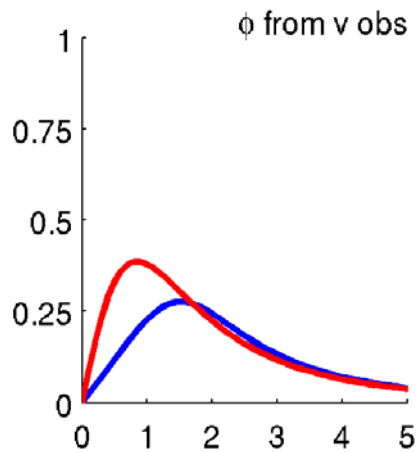
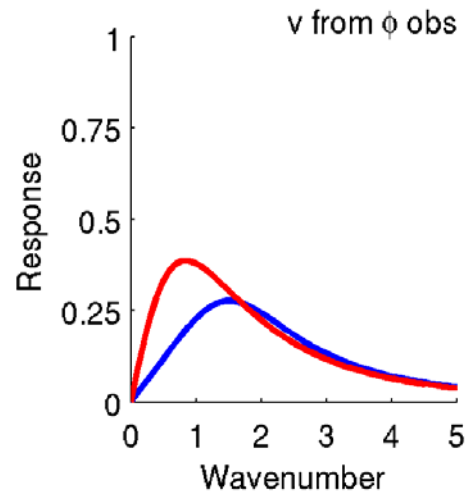
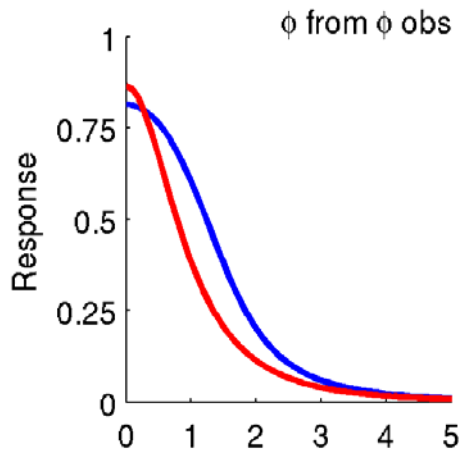
- Balance defined by null space of \mathbf{P}^f (Lorenc 2003)
- Localisation increases the rank of \mathbf{P}^f
 - necessary to fit observations
- Does localisation reduce the null space in a way that is consistent with balance?

Effect of localisation on geostrophic balance



- Localisation increases gradient of geopotential
 - Increases geostrophic wind
- Localisation reduces analysed wind speed
- Thus localisation weakens geostrophic balance
- Similar arguments apply for other balance relations

Dense observing system



- 1-dimensional infinite domain, observe wind and geopotential
- Geostrophic covariances
- Red: Kalman gain in spectral space, no localisation
- Blue: with localisation

Better localisation

- Covariances involving streamfunction ψ and velocity potential χ are more isotropic than those involving (u, v)
 - Long assimilation experience
- Balance equations relate $\text{grad}(\phi)$ to $\text{grad}(\psi)$
 - e.g. geostrophy, nonlinear balance equation
- Localising in (ψ, χ) rather than (u, v) space will be less severe on balance, because $\text{grad}(\phi)$ and $\text{grad}(\psi)$ will increase by similar amounts and because covariances in (ψ, χ) are more isotropic than in (u, v) .

Localisation in (φ, ψ, χ) -space

Write the velocity components in terms of streamfunction ψ and velocity potential χ :

$$\begin{aligned}u &= -\frac{\partial\psi}{\partial y} + \frac{\partial\chi}{\partial x} \\v &= \frac{\partial\psi}{\partial x} + \frac{\partial\chi}{\partial y}\end{aligned}\tag{1}$$

Write the two-point covariance

$$\langle u_1, u_2 \rangle = \langle u(x_1, y_1), u(x_2, y_2) \rangle\tag{2}$$

in terms of ψ and χ :

$$\begin{aligned}\langle u_1, u_2 \rangle &= \frac{\partial}{\partial y_1} \frac{\partial}{\partial y_2} \langle \psi_1, \psi_2 \rangle - \frac{\partial}{\partial y_1} \frac{\partial}{\partial x_2} \langle \psi_1, \chi_2 \rangle \\&\quad - \frac{\partial}{\partial x_1} \frac{\partial}{\partial y_2} \langle \chi_1, \psi_2 \rangle + \frac{\partial}{\partial x_1} \frac{\partial}{\partial x_2} \langle \chi_1, \chi_2 \rangle\end{aligned}\tag{3}$$

(φ, ψ, χ) -localisation (cont'd)

Let $C = C(x_1, y_1, x_2, y_2)$ be a localisation function, and apply the localisation by replacing $\langle \psi_1, \psi_2 \rangle \leftarrow C \langle \psi_1, \psi_2 \rangle$ in (3) and similarly for $\langle \psi_1, \chi_2 \rangle$, $\langle \chi_1, \psi_2 \rangle$ and $\langle \chi_1, \chi_2 \rangle$. Then the *localised* two-point covariance $\langle u_1, u_2 \rangle_L$ becomes

$$\begin{aligned} \langle u_1, u_2 \rangle_L &= C \langle u_1, u_2 \rangle \\ &+ \frac{\partial C}{\partial x_1} \langle \chi_1, u_2 \rangle - \frac{\partial C}{\partial y_1} \langle \psi_1, u_2 \rangle + \frac{\partial C}{\partial x_2} \langle u_1, \chi_2 \rangle - \frac{\partial C}{\partial y_2} \langle u_1, \psi_2 \rangle \\ &+ \frac{\partial^2 C}{\partial x_1 \partial x_2} \langle \chi_1, \chi_2 \rangle - \frac{\partial^2 C}{\partial x_1 \partial y_2} \langle \chi_1, \psi_2 \rangle - \frac{\partial^2 C}{\partial y_1 \partial x_2} \langle \psi_1, \chi_2 \rangle + \frac{\partial^2 C}{\partial y_1 \partial y_2} \langle \psi_1, \psi_2 \rangle \end{aligned} \quad (4)$$

The covariances in ψ , χ , u and v can all be calculated from the ensemble in the usual way. (The EnKF will need to read in ψ and χ in addition to u and v .)

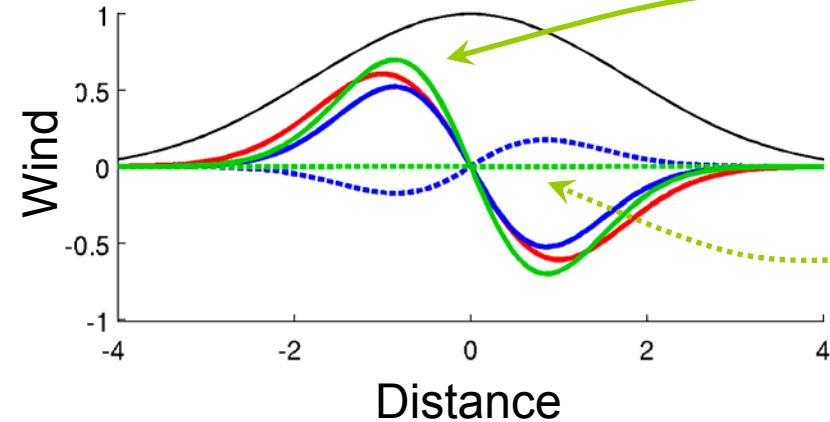
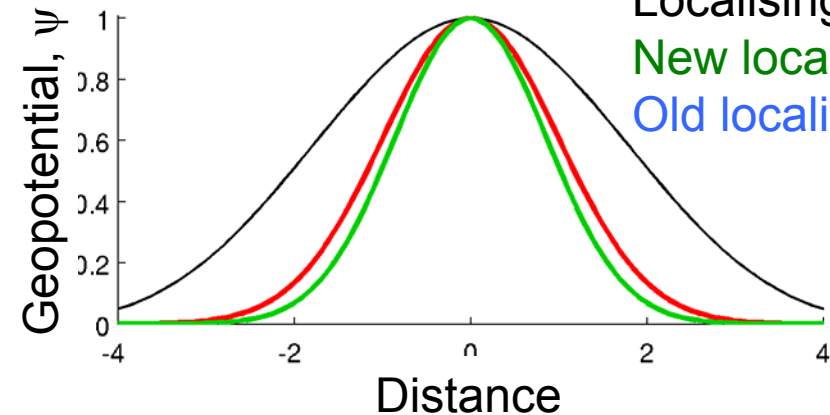
Other covariances are calculated similarly (Kepert, 2008).

Optional *intervariable localisation* with $\langle \psi_1, \chi_2 \rangle = \langle \phi_1, \chi_2 \rangle = 0$, etc.

Effect of localisation in (ϕ, ψ, χ) -space

Single height observation

True covariance (red)
Localising function (black)
New localisation (green)
Old localisation (blue)



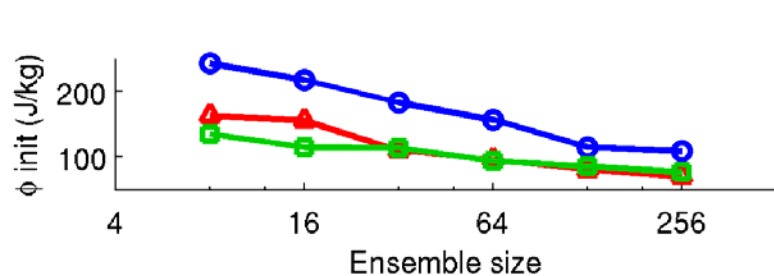
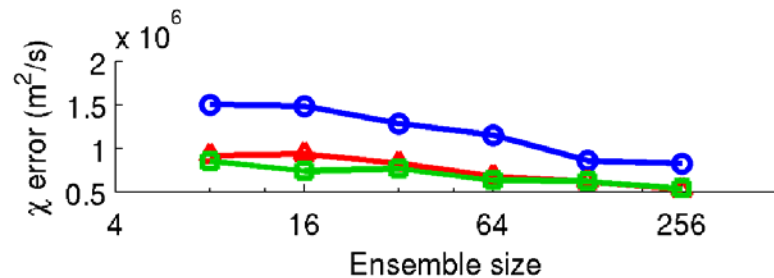
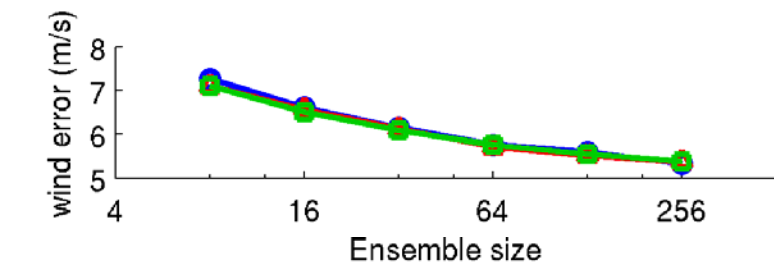
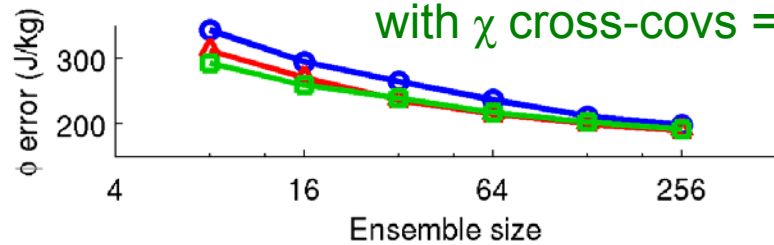
- Geopotential is same in both localisations.
- Wind response increases (because the ψ gradient increases)
- Ageostrophic wind in analysis is now zero.

Performance in test system

old localisation (blue)

new localisation (red)

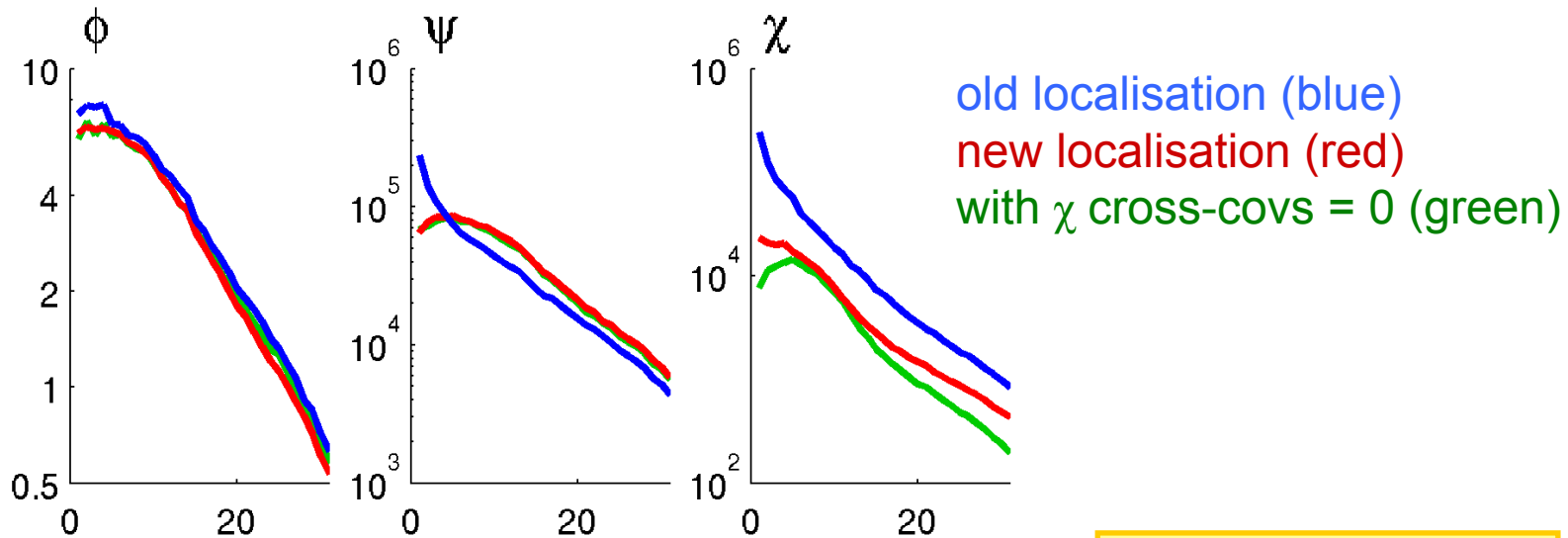
with χ cross-covs = 0 (green)



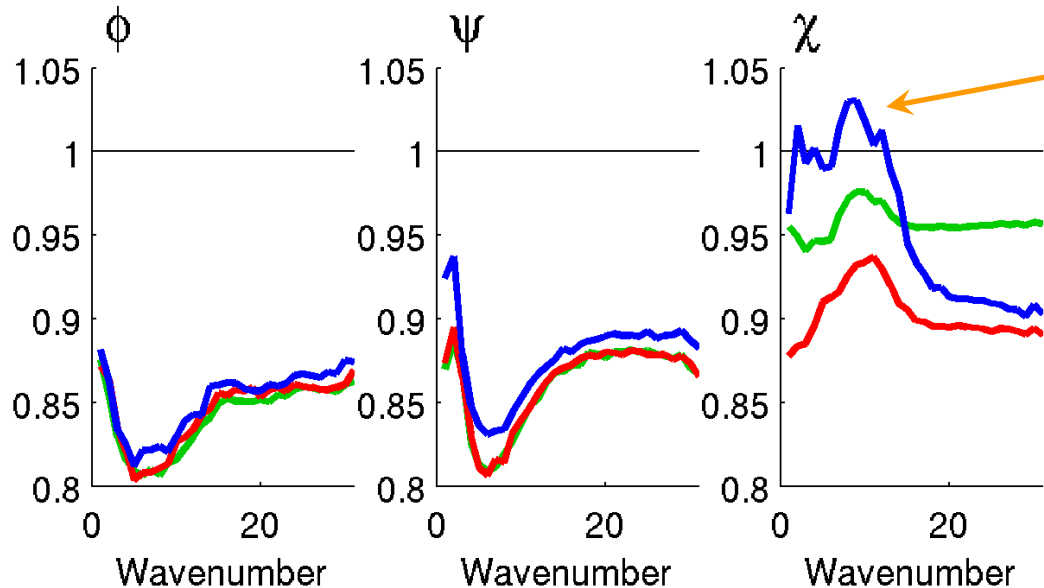
- Identical twin experiments, global spectral shallow-water model, various localisations.
- New localisations are substantially more accurate
- and better balanced.
- But analyses still contain some imbalance, and performance is improved when normal-mode initialisation is included (in this low-dissipation system).

Analysis increment spectra

increment in ensemble mean



relative increment in ensemble variance



Analysis increases ens. variance \rightarrow suboptimal gain (old localisation).

Less variance reduction when χ cross-covs = 0 in new localisations.

Observing mixed variables

- Most real observations are a mix of vortical and gravity modes
 - e.g. wind, geopotential, temperature, ...
 - assimilation has to correctly unscramble the mix
- Time-scale of vortical vs. gravity is variable
 - e.g. mid-latitudes vs. tropics, mesoscale
- Background variance of vortical vs. gravity modes is variable
 - e.g. troposphere vs. mesosphere
- Assimilation system needs to be robust in all these circumstances.

Idealised experiments with mixed-variable observations (Neef et al.)

- Accurate recovery of both modes requires
 - accurate fast-slow correlations
 - accurate error variances
- EnKF with enough members can do it
 - but tends to diverge in fast mode
 - and spurious projection onto fast mode can cause both to diverge
- Model error in gravity waves can be detrimental to EnKF performance
 - issue with semi-implicit time stepping schemes?
- Neef et al. (2006, 2008)

Observing mixed variables

Two-variable system, background covariance matrix

$$\mathbf{P}^b = \begin{bmatrix} \sigma_1^2 & \rho\sigma_1\sigma_2 \\ \rho\sigma_1\sigma_2 & \sigma_2^2 \end{bmatrix}$$

Observe the sum $\mathbf{H} = [11]$ with error $\mathbf{R} = [\sigma_o^2]$. The analysis covariance is

$$\mathbf{P}^a = \frac{1}{\sigma_1^2 + 2\rho\sigma_1\sigma_2 + \sigma_2^2 + \sigma_o^2} \begin{bmatrix} \sigma_1^2(\sigma_2^2(1 - \rho^2) + \sigma_o^2) & \sigma_1\sigma_2(\sigma_1\sigma_2(1 - \rho^2) + \rho\sigma_o^2) \\ \sigma_1\sigma_2(\sigma_1\sigma_2(1 - \rho^2) + \rho\sigma_o^2) & \sigma_2^2(\sigma_1^2(1 - \rho^2) + \sigma_o^2) \end{bmatrix}$$

so the variance ratio is

$$\frac{\sigma_1^2(\sigma_2^2(1 - \rho^2) + \sigma_o^2)}{\sigma_2^2(\sigma_1^2(1 - \rho^2) + \sigma_o^2)} = \frac{\frac{\sigma_1^2\sigma_2^2}{\sigma_o^2}(1 - \rho^2) + \sigma_1^2}{\frac{\sigma_1^2\sigma_2^2}{\sigma_o^2}(1 - \rho^2) + \sigma_2^2}$$

which is closer to 1 than σ_1^2/σ_2^2 .

If ρ^2 is too large due to sampling error, then the variable with least variance (e.g. gravity waves) will become underspread and may diverge.

Improving balance in an EnKF

- Keep everything as linear as possible
 - small analysis increments, small model error term, small covariance inflation, etc
 - eliminate or account for model bias (in all modes)
 - *requires an accurate system*
- Localise more carefully
- (Balanced) model error term
 - Better introduced at beginning of forecast?
- Initialise
 - Only way to stop the positive feedback?
- Generally less options than in VAR.

Bugbears

- Balance often not reported in literature
 - operational Canadian EnKF is an honourable exception.
- Relevant model settings (e.g. diffusion) sometimes not reported.
- Data voids are a good way of testing that covariances are ok (but not always done).
- Understanding EnKF balance properties may require different systems to practical applications (e.g. well-chosen toy models, very low dissipation, no initialisation).

Summary

- Balance is a significant issue in atmospheric EnKFs
- Multiple causes of imbalance, hard to diagnose in a full system. (Houtekamer and Mitchell 2005)
- Inherent to algorithm
 - because balances are nonlinear
- Various approximations also contribute
 - sampling error, localisation, model error, etc
- Some unbalanced flows are important and need to be captured
 - mesoscale, tropics, mesosphere
 - a global atmospheric system has to manage all these situations at once.